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ParisTech Intermittency in water cycle and geophysics: why so important?



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Nobel 2021 : complexity and intermittency

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to our understanding of complex physical systems"



Three different approaches, but a common object: our planet!





Emphasis of the prize scientific background on intermittency (*)

(*) Schertzer and Nicolis, Eos 2022)







Cascades as a generic case of complexity

Complex systems

- Complex system?
 - French Roadmap for Complex Systems: "A system comprised of a great number of heterogeneous entities, among which local interactions create multiple levels of collective structure and organization...".
 - explain failures to cope with their complexity at odds with reductionism (M. Gell-Man)
 - necessity to have a multiscale approach.









Rainfall

98.5

29

Forest (40.1)

1.3 + 2.4

Lake (2.7)

Soil

moisture

17



The terrestrial water balance does not include Antarctica

Global hydrological cycle (Oki & Kanae , 2006)

Water cycleS













Art piece 'Windswept' (Ch. Sowers, 2012): 612 freely rotating wind direction indicators to help a large public to understand the complexity of environment near the Earth surface

Intermittency



Polarimetric radar observations of heavy rainfalls over Paris region during 2016 spring (250 m resolution):

- heaviest rain cells are much smaller than moderate ones
- complex dynamics of their aggregation into a large front



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Purely random landscape, but highly structured, (S+L, 1992)

Intermittency

- 1950s
- now a major transdisciplinary subject unusual features
- - extreme concentration of activity
 - structuring fluctuations
 - a multiplicity of anomalous scaling laws
 - (multifractal) phase transitions for extremes
- recent understanding of their interrelationships, • Legendre transform between concentration on a hierarchy of supports and the multiplicity of scaling laws
- - but also unresolved debates



• a matter for a few turbulence specialists in the



(Billion)



Growing Urban Population



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... in the context of climate change

- How to achieve resilient cities and mobility? Which strategies of reduction/adaptation to climate change? How to downscale climate scenarios to relevant scales?



Strong agreement of models on temperature increases...





but strong **disagreement** on the evolution of precipitation extremes





SPM 2.2 Projected changes in the climate system

Surface temperature is projected to rise over the 21st century under all assessed emission scenarios. It is very likely that heat waves will occur more often and last longer, and that extreme precipitation events will become more intense and frequent in many regions. The ocean will continue to warm and acidify, and global mean sea level to rise. {2.2}



Dotted areas: projected change larger than natural internal variability

Hatched areas: projected change less than one standard deviation than natural internal variability

Figure SPM.7 | Change in average surface temperature (a) and change in average precipitation (**b**) based on multi-model mean projections for 2081–2100 relative to 1986–2005 under the RCP2.6 (left) and RCP8.5 (right) scenarios.





With every increment of global warming, regional changes in mean climate and extremes become more widespread and pronounced



IPCC AR6







Louis Navier (1822)

Augustin-Louis Cauchy

Adhémar Jean Claude Barré de Saint-Venant

The fundamental equation of fluid dynamics derived at Ecole Nationale des Ponts et Chaussées

Two centuries ... since Navier (1822)







Simeon Denis Poisson

Sir George Stokes (1843)









Lewis Fry Richardson's "dream" (1922)

WEATHER PREDICTION BY NUMERICAL PROCESS

LEWIS F. RICHARDSON, B.A., F.R.MET.Soc., F.INST.P.

FORMERLY SUPERINTENDENT OF ENGLALEMULE OBSERVATORY LECTURER ON PHYNCS AT WEFTMINITER TRAINING COLLEGE



64.000 « human computers» synchronised by a « coordinator »

(François Schuiten, 2000)





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Modelling with brute force?

- Multiscale complexity vs. computing brute force
 - –(nonlinear) equations discretised on N voxels (≈cubes)
 - # voxels (1mm³⁾ to reach the viscous scale (≈1mm):
 - atmosphere ≈ 10 km high x (10,000 km)² horizontal
 - => $N \approx 10^7 (10^{10})^2 = 10^{27}$
 - much larger than $N_A = 10^{23}$!!
- Why to be deterministic... over only a small range of scales?



!!
er only a





Louis Navier (1822)

Augustin-Louis Cauchy





A millenium problem raised at Ecole Nationale des Ponts et Chaussées, recent episodes: Otelbaev (2013) and Tao (2015)

Two centuries ... since Navier (1822)





Adhémar Jean Claude Barré de Saint-Venant

Simeon Denis Poisson

Sir George Stokes (1843)











WEATHER PREDICTION BY NUMERICAL PROCESS

and the second

BY

LEWIS F. RICHARDSON, B.A., F.R.MET.Soc., F.INST.P.

PORMERLY SUPERINTENDENT OF EIKDALEMUIR OBSERVATOR LECTURER ON PHTNES AT WEITMINITER TRAINING COLLEG the sketch can be completed. We realize thus that: big whirls have little whirls that feed on their velocity, and little whirls have lesser whirls and so on to viscosityin the molecular sense.



A century of cascades !



The Supply of Energy from and to Atmospheric Eddies. By LEWIS F. RICHARDSON.

(Communicated by Sir Napier Shaw, F.R.S. Received March 9, 1920.)

Introduction.

This paper extends Osborne Reynold's theory of the Criterion of Turbulence, to make it apply to the case in which work is done by the eddies, acting as thermodynamic engines in a gravitating atmosphere. For







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Scale symmetry and equations

solutions are still unknown (existence, uniqueness): $\frac{\partial \underline{u}}{\underline{u}}$ $+\underline{u}$.

one can point out a scale symmetry (*):

$$\tilde{T}_{\lambda}: \quad \underline{f} \mapsto \underline{f}$$

Kolmogorov's scaling (K41 obtained with:

(*) to generalised Galilean invariance (S+al, 2010)



Whereas elementary mathematical properties of Navier-Stokes

$$\underline{grad}(\underline{u}) = \underline{f} - \frac{1}{\rho} \underline{grad}(p) + \nu \Delta \underline{u}$$

 $T_{\lambda} \colon \underline{x} \mapsto \underline{x}/\lambda$

$$t \mapsto t/\lambda^{1-\gamma}; \underline{u} \mapsto \underline{u}/\lambda^{\gamma};$$

$$\lambda^{2\gamma-1}; \nu \mapsto \nu/\lambda^{1+\gamma}; \rho \mapsto \rho/\lambda^{\gamma}$$

1)
$$\varepsilon(\ell) \approx \frac{\delta u^3}{\ell} \approx \bar{\varepsilon} \Rightarrow \gamma_K = 1/3$$

- General case: multiple singularities γ 's: $Pr(\gamma' > \gamma) \approx \lambda^{-c(\gamma)}$
 - from to self-similarity (Sedov, 1961), symmetry (Parisi +Frisch, 1985),



Symmetry in the Fourier space

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Kolmogorov-Obukhov spectrum (K41)



$$\underline{x} \mapsto \underline{x}/\lambda; \ \underline{u} \mapsto \underline{u}/\lambda^{\gamma}$$

$$\underline{\widehat{u}}(\underline{k}) = \int_{-\infty}^{\infty} d\underline{x} \exp(i\underline{k}.\underline{x}) \ \underline{u}(\underline{x})$$

$$\underline{k} \mapsto \underline{k}\lambda; \ \underline{\widehat{u}} \mapsto \underline{\widehat{u}}/\lambda^{\widehat{\gamma}}$$

$$\Rightarrow \widehat{\gamma} = d + \gamma$$
extral scaling: $\langle u^2 \rangle = \int_{0}^{\infty} dk \ E$

$$E(k) \propto k^{-2\gamma_{K}-1} = k^{-5/3}$$





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Scale symmetry and equations

solutions are still unknown (existence, uniqueness):

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Kolmogorov's scaling (K41 obtained with:

General case: multiple singularities γ 's:

to generalised Galilean invariance (S+al, 2010)



- Whereas elementary mathematical properties of Navier-Stokes
 - $\frac{\partial \underline{u}}{\partial t} + \underline{u} \cdot \underline{grad}(\underline{u}) = \underline{f} \frac{1}{\rho} \underline{grad}(p) + \nu \Delta \underline{u}$
 - $T_{\lambda} \colon \underline{x} \mapsto \underline{x}/\lambda$

$$t \mapsto t/\lambda^{1-\gamma}; \underline{u} \mapsto \underline{u}/\lambda^{\gamma};$$

$$/\lambda^{2\gamma-1}; \nu \mapsto \nu/\lambda^{1+\gamma}; \rho \mapsto \rho/\lambda^{\gamma'}$$

1)
$$\varepsilon(\ell) \approx \frac{\delta u^3}{\ell} \approx \bar{\varepsilon} \Rightarrow \gamma_K = 1/3$$

$$\Pr(\gamma' > \gamma) \approx \lambda^{-c(\gamma)}$$

from to self-similarity (Sedov, 1961), symmetry (Parisi +Frisch, 1985),

From scaling analysis to scalar cascade processes 10"









multiplication by 4 independent random (multiplicative) increments

- - . . .
 - - ß-model: Novikov and Stewart (1964), Frisch et al. (1978)
 - Log normal model: Yaglom (1966)
 - Limit log-normal: Mandelbrot (1974)
 - $-\alpha$ -model: S+L (1984, 1985)
 - Multiplicative chaos: Kahane (1985)
 - Universal multifractals/Levy multiplicative chaos: S+L (1987a&,b, 1997)), Fan (1987)

 - Log-Poison model: Dubrulle (1994)

Schertzer & Lovejoy, 1989b

— . . .

Scaling analysis

- Passive scalar dispersion: Richardson (1926)
- Structure function: Kolmogorov (1941)
- Energy spectrum: Obukhov (1941)
- Higher order structure functions: Kolmogorov (1962), Obukhov (1962) – Renyi dimensions: Grassberger and Procaccia (1983), Hentschel, and Procaccia (1983)
- Legendre transform to dimensions: Parisi and Frisch (1985)
- Fractal measures: Halsey et al. (1987)

Cascade processes



Non-intermittent ("homogeneous") cascade

 $N(\ell_n) \approx \ell_n^{-D}$ L/2 L/2

β-model





 β -model of intermittency : sub-eddies are either "alive" or "dead"







Varenna summer school (1983)



Fig.1 S+L (1984) fitted the α -model to the experimental data of Anselmet et al (1983)

• it was shown to fit the experimental points from Anselmet et al. (1983), see fig. 1 with $H = 1/3, \alpha = 5, 5.5, 6$



- "Turbulence and Predictability in Geophysical Fluid" Dynamics" organised by M. Ghil, R. Benzi et G. Parisi
 - -a primary version of the multifractal formalism of Parisi and Frisch (1985) that concluded: "Still the multifractal model appears to be somewhat more restrictive than Mandelbrot's weightedcurdling model which does include the logornormal case".
 - S+L's short presentation (Varenna proceedings (1985)) refers to S+L (1984)):
 - a small perturbation of the β -model is no longer limited to a unique dimension (α -model)
 - the divergence of higher order moments is rather generic
 - it introduces spurious scaling, an analytical approximation depending on a unique scaling exponent H and the critical order α was proposed:

$$\xi(p) = pH + \theta(p-\alpha)(1-p/\alpha)$$







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 Example of Kolmogorov's refined scaling laws (K62, $0 \le q \le q_D$):

Structure function of order *q*

Normal/dimensional scaling (K41) unique law (linear exponent wrt q)

$$\delta_r u := u(x+r) - u(x)$$



$S_q(r) := \langle \delta_r u^q \rangle \propto \bar{\varepsilon}^{q/3} r^{q/3} (r/L)^{C_1(q/3-1)q/3}$

Anomalous/adimensional scaling multiple laws (nonlinear exponent wrt q)

refined scaling (K62)





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Measure of the **Fraction** at resolution λ $(\ell = L/\lambda)$ of the **Probability Space** corresponding to the event $\{\epsilon_{\lambda} \geq \lambda^{\gamma} \overline{\epsilon}\}$:

 $\Pr(\varepsilon_{\lambda} \geq \lambda^{\gamma} \overline{\varepsilon}) \approx \lambda^{-c(\gamma)}$

 $c(\gamma)$ is a codimension, γ a singularity

(Schertzer and Lovejoy, 1987)

or a 'Cramer's function' (Mandelbrot, 1991)

Multifractal Fields





high singularities

medium singularities

low singularities



Cascades and statistical physics

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$$\varepsilon_{\lambda} = e^{\Gamma_{\lambda}}; \quad Ee^{q\Gamma_{\lambda}} = Z_{\lambda}(q) = e^{K_{\lambda}(q)} \approx \lambda^{K(q)} \longleftrightarrow P(\varepsilon_{\lambda} > \lambda^{\gamma}) \approx \lambda^{-c(\gamma)}$$

$$\downarrow$$
Legendre
ce of the generator
$$K(q) \longleftrightarrow c(\gamma)$$
les Statistical physics analogues
$$K(q) = \max[q\gamma - c(\gamma)]; \quad c(\gamma) = \max[q\gamma - c(\gamma)];$$

main "trick": log-divergend

> Cascad Γ : generator \approx hamiltonian q: statistical order \approx inverse of temperature Z: 1st characteristic or moment generating function \approx partition function K: 2nd characteristic or cumulant generating function \approx Gibbs free energy c: codimension or Kramer function \approx entropy γ : singularity or Hölder exponent \approx energy



- Transformation of a measure σ with the help of a "density" ε into another measure Π : $d\Pi = \varepsilon d\sigma$
- Generalisation with a non trivial limit ε of densities ε_{λ} of increasing resolution: $\lambda = L/\ell \to \infty$

$$K(q) = \max_{\gamma} [q\gamma - c(\gamma)]; \ c(\gamma) = \max_{q} [q\gamma - q\gamma]$$

-K(q)



To downscale climate scenarios?



- Consensus on the need to downscale simulations: to manage water
- below a kilometer to model cities ? but a fondamental obstacle :
 - relationships large/small scales
 - stationnary? (« stationnarity is dead ! »)
 - Implicit hypothesis in:
 - correlations (linear or not) large/small scales or neural networks: training on « non perturbated » data sets; - weather types: stationnary basis, only frequences change; - GCM: stationnary parametrisationsof small scales...

Alternative:

- Analyze and take into account the evolution of fluxes through scales!
- Example: **rainrate** for the water cycle. \bullet







1. Parsimonious description: only 3 exponents:

- intermittency:
 - average intermittency C_1 : how sparse is the average rainfall?

 $C_1 \neq 0$: it does not rain everyday, everywhere!

- intermittency variability α : diversity of rainfall regimes
- $\alpha \neq 0$: not only the alternative rain/no rain!
- scale dependence **H** of the average rainfall $\langle R_1 \rangle$? For rainfall **H**≈0
- **Trivial conséquences for the extremes:** a

More generally: a

- allow to compare quantiles over various durations

Scaling of precipitations

- C_1 and $\alpha \uparrow =>$ extremes \uparrow or C_1 and $\alpha \downarrow =>$ extremes \downarrow
- these exponents define the Intensity-Duration-Frequency (IDF) curves over a wide range of scale and intensities (Benjdoudi et al., 1987, Tchiguirinskaia et al 2010);



Analysis of the scenario A2

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0.48

0.47

0.46

1860

04.01

 1980

1900



2020

2060

31

29

27



average intermittency C1 ↑
intermittency
variability α ↓,
> difficulty to
evaluate extremes of
precipitations



=> refined analysis :

- time volution of the Most Probable Singularity

γ_s (Hubert et al, 1993; Douglas & Barros, 2003):

- a scale invariant statistic, more stable than

the maximal simulated precipitation P_{max}.

-Unable us to conclude: **extremes (Royer et al., 2008)**,

-- seasonality can be taken into account (Royer et al., 2010)



Mapping H



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Hurst parameter H: scaling exponent of the mean field (*H* = 0 for conservative multiplicative cascade):

for Canada (~equally 25%): 0.6-0.25 (red); 0.25-0.2 (yellow); 0.2-0.17 (green); 0.01-0.17 (blue)

and for US: 0.6-0.25 (~ 19%, red); 0.25-0.2 (~ 19%, yellow); 0.2-0.17 (~ 14%, green); 0.01-0.17 (~ 48%, blue)



Mapping C1



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Codimension C1: how intermittent is the mean field (~equally 25%): 0.9-0.2 (red); 0.15-0.2 (yellow); 0.1-0.15 (green); 0.01-0.1 (blue)







Mapping a

Multifractality index α how multifractal is the process

>2 (~ 1%; black); 2.0-1.5 (~ 30%, red); 1.5-1.0 (~ 52%, yellow); 1.0-0.6 (~ 15%, green); <0.6 (~2%, blue)

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Histograms of multifractal parameters

- 3. Main results
- 3.1 Database
- 3.2 Analysis

Parameter distributions for empirical data \rightarrow

Multifractal parameter distribution

'prior' probabilities Pr(CL,C1)

Multifractal Flood Frequency analysis: quantile plots and codimension

 α , C1, H \rightarrow critical singularity γ_{D}

α

Predict shape of multifractal codimension function:

$$C(\gamma) = C_1 \left[\frac{1}{\alpha} \left(\frac{(\alpha - 1)}{C_1} \gamma + 1 \frac{1}{\frac{1}{2}} \right]^{\alpha - 1} \right]$$

Transform into probability space:

$$T = \lambda^{c(\gamma)}$$

$$Q/\langle Q \rangle = \lambda^{\gamma - H}$$

Mapping the reference discharges: $Q(\gamma_D)$

Reference discharge

for Canada: >1000 (~14%, red); 300-1000 (~17%, yellow); 50-300 (~24%, green); (~45%, blue) <50

for US: >30000 (~14%, red); 8000-30000 (~22%, yellow); 1000-8000 (~38%, green); (~26%, blue) <1000

Mapping the reference discharges: T=100

Discharges corresponding to 100 year return period

for Canada: >1000 (~23%, red); 300-1000 (~16%, yellow); 50-300 (~26%, green); <50 (~35%, blue)

For US: >30000 (~18%, red); 8000-30000 (~27%, yellow); 1000-8000 (~36%, green); <1000 (~19%, blue)

Multifractal Flood Frequency analysis: Kenogami series

Red dots are empirical frequencies for data sample; Red line is "classical" approach; Blue dots are MMFA results;

1996 Saguenay flood of 2000 m³/s has return period of:

-1000 years for UM,

-billions of yeas for classical analysis

Testing acceptability criteria on annual maxima: sample size

These parameter fluctuations for Kenogami series give quite a bit of dispersion for high return periods

→ uncertainty/reliability ?

Scaling & multifractality of GRIP data

Figure 1. High resolution (200 yr average) $\delta^{18}O$ record coming from the recent GRIP Greenland ice core. Sharp fluctuations occuring on small time scales are clearly visible.

 $E(f) \approx f^{-\beta}; \ \beta \approx 1.4$

СТО ГОДИНА МИЛАНКОВИЋЕВЕ ТЕОРИЈЕ О КЛИМАТСКИМ ПРОМЕНАМА

100 YEARS of **MILANKOVIC'S THEORY of CLIMATE CHANGE**

Figure 2. The power spectrum of the $\delta^{18}O$ data, shown in a log-log plot. The global straight line $f^{-1.4}$ is an indication of scaling. There is no evidence of periodic variations of frequencies of $(20 \text{ kyr})^{-1}$ or $(40 \text{ kyr})^{-1}$.

(Schmitt et al., GRL, 1995)

Scaling & multifractality of GRIP data

Figure 3. The structure functions $<(\Delta \theta_{\tau})^q > vs. \tau$ in a log-log plot for q=1, 1.5, 2, 2.5, and 3. Linear trends are clearly visible for all order of moments, for a range of scales going from ≈ 400 yr to ≈ 40 kyr. The straight lines indicate the best linear regression over this range of scales for each value of q. This gives in particular: $H = \zeta(1) = 0.24 \pm 0.02$, and $\zeta(2) = 0.40 \pm 0.03$.

$$< \left(\Delta \theta \left(\frac{T}{\lambda}\right)^q >$$

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Figure 4. The scaling exponent structure function $\zeta(q)$ empirical curve (dots), compared to the monofractal curve $\zeta(q) = qH$ (dashed and dotted line), and to the universal multifractal function obtained with $\alpha = 1.6$ and $C_1 = 0.05$ in equation (2) (dotted curve). The empirical curve is nonlinear, indicating multifractality. For moments of order q>1, the departure from monofractal behaviour is strong. The universal multifractal fit is excellent until moment order $q_{\text{max}} \approx 2.5 \pm 0.3$, corresponding to a multifractal phase transition occuring because of sample limitations.

 $> \approx \lambda^{-\zeta(q)} < (\Delta \theta(T)^q > 1)$

(Schmitt et al., GRL, 1995)

Scaling & multifractality of GRIP data

Figure 5. Empirical curve of $\log[q\zeta(0) - \zeta(q)]$ vs. $\log(q)$ giving directly the universal multifractal indices α and C_1 : α is the slope of the straight line and C_1 is estimated by the intercept. We obtain $\alpha \approx 1.6$ and $C_1 \approx 0.05$.

СТО ГОДИНА МИЛАНКОВИЋЕВЕ ТЕОРИЈЕ О КЛИМАТСКИМ ПРОМЕНАМА

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linear scaling, fractional integration

nonlinear scaling intermittency, multifractality

 $\beta = 2H + 1 + K(2) \approx 2H + 1$

H (fractional) order of integration C_1 codimension of the mean intermittency α multifractality index, concavity of $\zeta(q)$

 $H \approx 0.2, C_1 \approx 0.05, \alpha \approx 1.6$

(Schmitt et al., GRL, 1995)

Vostok data

Age (yr BP)

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A very rich data set over 450 000 years!

(Petit et al., Nature, 1999)

Vostok data

Figure 4 Spectral properties of the Vostok time series. Frequency distribution (in cycles yr⁻¹) of the normalized variance power spectrum (arbitrary units). Spectral analysis was done using the Blackman-Tukey method (calculations were performed with the Analyseries software⁴⁷): **a**, isotopic temperature; **b**, dust; **c**, sodium; **d**, $\delta^{18}O_{atm}$; **e**, CO₂; and **f**, CH₄. Vertical lines correspond to periodicities of 100, 41, 23 and 19 kyr.

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"integrated spectra": $\omega E(\omega) vs \cdot Log(\omega)$

well identified but broad peaks/spikes

(Petit et al., Nature, 1999)

Vostok data: glaciological timescale <u>increments</u>

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(S&T, 2020)

time (yr)

Vostok data: spectral scaling

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Vostok data: multi-scaling

Structure functions ≈ generalisation of Fourier spectra ($q \neq 2$):

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$$< \left(\Delta \theta \left(\frac{T}{\lambda}\right)^q > \approx \lambda^{-\zeta(q)}$$

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initial rotation along the smallest axis

Gyroscopic instability

initial rotation along the medium axis

wobble & gimbal lock

Hyperion

Apollo 13 guidance computer console

3D Scaling Gyroscope Cascade

Local flux of energy:

$$\varepsilon_{n}^{i} = -\sum_{r=0}^{n} k_{n-r+1} \left[\left| \hat{u}_{n-r+1}^{2a^{r}(i)-1} \right|^{2} - \left| \hat{u}_{n-r+1}^{2a^{r}(i)} \right|^{2} \right] \operatorname{Im}(\hat{u}_{n}^{2a^{r}(i)})$$

 $\hat{u}_{n}^{a^{r}(i)}_{n}^{(i)+(-1)} = a^{r(i)+1} k_{n-r} \left| \hat{u}_{n}^{a^{r+1}(i)} \right|^{2} \operatorname{Im}(\hat{u}_{n-1}^{a^{r+1}(i)})$ T&S, T. al, 1998, Li et al. 2023

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Log K(q=1.5,¹)

energy flux).

SGC-3D has a 1st order phase transition: FAT TAILED distribution Shell model (SGC-1D) has not !

SGC vs. shell-models

Scaling of DTM curves with q=1.5, $\eta=\overline{0.5}$, 3.5 (for simulated

Comparison	Lévy index of the	codimension
between:	generator	of the mean
		fractality
SGC-3D space	$\alpha = 1.4 \pm 0.05$	$c_1 = 0.25 \pm 0.05$
SGC-3D time	$\alpha = 1.5 \pm 0.05$	$c_1 = 0.25 \pm 0.05$
Shell-model of the	$\alpha = 0.6 \pm 0.05$	$c_1 = 0.4 \pm 0.05$
SGC (time)		

Mr Jourdain and Lie cascades

What is general and theoretically straightforward:

scalar valued cascade: $R^d \rightarrow R^+$

- Lie group: smooth manifold
- Lie algebra: tangent space to the group at the identity
- therefore a vector space with a skew product that satisfy the Jacobi identity:

[X, [Y, Z]] + [Y, [Z, X]] + [Z, [X, Y]] = 0

exemple: commutator of matrices

• $exp: Lie algebra \mapsto Lie group$

x vM $\gamma(t)$

 $[X, Y] = XY - YX \qquad [X, Y] = 0 \Rightarrow \exp(X + Y) = \exp(X) \exp(Y)$

Surface layer complexity!

Art piece 'Windswept' (Ch. Sowers, 2012): 612 freely rotating wind direction indicators to help a large public to understand the complexity of environment near the Earth surface

Both movies illustrate the challenge of the near surface wind that plays a key role in the heterogeneity of the precipitations... and wind energy!

Multifractal FIF simulation (S et al., 2013) of a 2D+1 cut of wind and its vorticity (color). This stochastic model has only a few parameters that are physically meaningful.

Fractionnaly Integrated Flux model (FIF, vector version)

FIF assumes that both the renomalized propagator G_R and force f_R are known:

$$G_R^{-1} * u = f_R$$

where: $f_R = \varepsilon^a$

$$G_R^{-1}$$

is a fractionnal differential operator results from a continuous, vector, multiplicative cascade (Lie cascade) 3D FIF wind simulation based on quaternions

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- UN 2030 Agenda and follow-ups: a vibrant call to engineering intelligence
- not achievable without
 - -advanced observation technologies (e.g. small radars and lidars)
 - -innovations (e.g., Blue Green/Nature Based solutions)
 - -synergic, integrative approaches overcoming the traditional silo thinking
 - -disruptive methodologies (multiscale analysis, simulation and modelling) for full deployment of the above

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- Intermittency: a key issue in physics, mechanics, geophysics and elsewhere
- Breakthroughs with cascades and multifractals
 - intermittency across scales
 - New approaches of extremes, predictability and predictions
 - Beyond commonalities significant differences of formalisms
- Beyond scalar valued fields
 - Scaling Gyroscopes Cascades
 - Multifractal operators: exponentiation of a stochastic Lie algebra of generators
 - Multifractal multivariate analysis

S&T, Earth& Space, 2020 Chaos 2015, S&al. ACP, 2012 S&L, IJBC, 2011, Fitton&al., JMI 2013

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Final conclusion: Philip Anderson (Phys. Today 41 526) 1988):

"A real scientific mystery is worth pursuing to the ends of the Earth for its own sake, independently of any obvious practical importance or intellectual glamour."

Intermittency is without doubt such a mystery, but not without multifaceted practical importance and numerous corresponding contributions.

This is more than being illustrated by thousands of communications in EGS/EGU NP since 1988, in particular those of our colleagues of HM&Co.

