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1	Anisotropic thermal conductivity of natural Boom Clay
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31 Abstract

32 The thermal conductivity of host rocks is an important parameter in the design of deep geological 33 disposal of heat-emitting radioactive waste. Due to bedding, heat transfer in sedimentary rocks is 34 affected by their transversally isotropic structure. In this work, an experimental program is run to 35 measure the thermal conductivities of Boom Clay along various orientations with respect to the 36 bedding plane by using the needle thermal probe technique. Measurements were performed on 37 specimens obtained from cores drilled from the HADES Underground Research Laboratory (URL) at 38 Mol, Belgium, at a depth of 223 m. The thermal conductivities values obtained are in good agreement 39 with those previously published, confirming the thermal anisotropy of the Boom Clay. Moreover, the 40 observed changes in thermal conductivity with respect to the distance to the gallery provide further 41 evidence on the extent of the Excavation Damaged Zone around the gallery.

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43 Keywords: Thermal conductivity; Boom Clay; thermal anisotropy; excavation damaged zone;
44 laboratory test.

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56 1 Introduction

57 In the context of deep geological disposal of heat-emitting high activity radioactive waste, the thermal

58 conductivity of the host geological formation (stiff clays, mudstones, granites, etc.) is an important

59 parameter in the design of the disposal. Moreover, the thermal transient phase that will be experienced 60 by the system over hundreds to thousands of years after waste emplacement represents an important 61 part of its long-term evolution (Yu et al., 2013). Due to a layered microstructure resulting from the 62 initial deposition and further geological processes, stiff clays and claystones are known to 63 preferentially conduct heat along the direction parallel to the bedding plane. Indeed, this anisotropy of 64 thermal conductivity was observed in various stiff clays and claystones including the London clay, the 65 Callovo-Oxfordian claystone (COx, East of France) and the Opalinus clay (Switzerland). In London 66 clay, the thermal conductivity in the direction perpendicular to the bedding plane was found equal to 67 0.83 W/(m.K) whereas that in the direction parallel to the bedding plane was equal to 1.19 W/(m.K) 68 (Midttømme and Roaldset, 1999). The thermal conductivity of Opalinus Clay was investigated by 69 Mügler et al. (2006). They reported that the values obtained by in-situ back analysis are 0.55 - 1.07 70 W/(m.K) in the direction perpendicular to bedding and 1.84 - 1.90 W/(m.K) in the direction parallel to 71 bedding, respectively. Buntebarth (2004) (in Jobmann and Polster, 2007) gave, from laboratory 72 experiments, thermal conductivities of 0.75 W/(m.K) and 1.55 W/(m.K) for the directions 73 perpendicular and parallel to bedding, respectively. The thermal conductivity in the direction 74 perpendicular to the bedding plane of the COx claystone varies between 1.3 and 1.9 W/(m.K) whereas 75 that in the direction parallel to the bedding plane varies between 1.9 and 2.7 W/(m.K) (ANDRA, 76 2005).

Based on the values in the two directions, a degree of thermal conductivity anisotropy (or anisotropy effect) η can be defined, as follows (Popov et al., 1999; Pribnow et al., 2000; Davis et al., 2007):

$$\eta = \frac{\lambda_{\prime\prime}}{\lambda_{\perp}} \tag{1}$$

From the values of thermal conductivity introduced above, a value of $\eta = 1.43$ is determined for London clay. For Opalinus Clay and COx claystone, η can vary from 1.78 to 3.34 and from 1.42 to 1.46, respectively.

In Boom Clay, the anisotropy of thermal conductivity has been inferred through the back analysis of a small scale in-situ heating experiment, ATLAS III, carried out in the Underground Research Laboratory (URL) HADES at a depth of 223 m (Chen et al., 2011). The measurements of temperature evolution at different distances from the heater were compared with the results of heat transfer 86 simulations. In those simulations, adopting values of thermal conductivities equal to 1.65 W/(m.K)87 and 1.31 W/(m.K) for the directions parallel and perpendicular to the bedding plane, respectively, led 88 to good agreement with the measured temperatures. It is worth noting that, in some sedimentary 89 rocks, the thermal conductivity in the direction parallel to the bedding plane is more than twice higher 90 than that in the direction perpendicular to the bedding plane (Midttømme et al., 1996). After Schön 91 (1996), the reasons for this anisotropy in thermal conductivity of sedimentary rocks are: (i) crystal 92 anisotropy of the individual rock-forming minerals; (ii) intrinsic or structural anisotropy resulting from 93 the mineral shapes and their textural arrangement within the rock; (iii) orientation and geometry of 94 crack fractures and other defects. Point (ii) is particularly true in the case of clays and claystones 95 because of the planar shape of clay platelets and the bedding orientation by long term consolidation, 96 creep and diagenesis.

97

98 In addition to anisotropy, the presence of the Excavation Damage Zone (EDZ) around underground 99 constructions can also influence the thermal conductivity. However, few studies have been devoted to 100 this aspect.

101

102 Three methods are often used to measure the thermal conductivity of soils in the laboratory: i) the 103 divided bar method, ii) the needle probe method (or line source method) and iii) optical scanning. The 104 divided bar method is a steady state method in which a constant temperature gradient is imposed 105 across the sample, resulting in a stable heat flow. This method involves heat transmission along a 106 single direction only, and is considered to be accurate (Johansen, 1975; Farouki, 1981; Brigaud et al., 107 1990; Midttømme and Roaldset, 1999). It is also recommended to be used in measuring the thermal 108 conductivity of rock samples (Beck, 1988; Ashwroth, 1990). The needle probe method is a transient 109 method in which a radial heat flow is produced within the specimen while measuring temperature 110 changes over time. This method is based on the theory of axisymmetric heat diffusion from an infinite 111 line source within an infinite surrounding medium. It is valid when the ratio between the length and 112 the diameter of the needle is larger than 30:1 (Popov et al., 1999). The thermal conductivity can hence 113 be back calculated from the temperature-time curve (Jessop, 1990). In the optical method (Popov, 114 1983), a focused, mobile and continuously operated heat source is used to heat a sample while

115 monitoring temperature changes using infrared temperature sensors. The excess temperatures 116 monitored on the sample are compared to that of a reference sample of known thermal conductivity, 117 allowing the thermal conductivity to be determined. This method has been recommended to estimate 118 the thermal anisotropy of rocks (Popov et al., 1999).

119

120 In this paper, the needle probe method was selected for its simplicity and standardisation. The thermal 121 anisotropy of Boom Clay was investigated by adopting various orientations of the needle with respect 122 to the bedding orientation of the samples. These experimental results were then compared with the 123 theoretical results. The Boom Clay samples were extracted from a series of cores taken from a 124 horizontal borehole drilled from a gallery of the HADES URL. The core was at 2.40 m to 22.3 m 125 distance to the gallery axis (HADES borehole 2012-2 / Connecting Gallery / Ring 66-67W) and the 126 specimens were extracted from the core at various distances ranging from 2.5 to 20.8 m to the gallery 127 axis.

128

129 **2 Background**

130 **2.1** Anisotropy of thermal conductivity

131 Heat flow by conduction along one dimension is governed by Fourier's law:

$$q = -\lambda \frac{dT}{dx}$$
(2)

where *q* is the heat flux density (W/m²), λ is the thermal conductivity (W/(m.K)) and *dT/dx* is the temperature gradient along the *x* direction.

In an anisotropic material, the apparent thermal conductivity λ measured on an anisotropic sample around a line-source is related to the principal components of the thermal conductivity tensor (Grubbe et al., 1983; Popov et al., 1999), as follows:

$$\lambda = \sqrt{\lambda_A \lambda_B \cos^2(\gamma) + \lambda_A \lambda_C \cos^2(\beta) + \lambda_B \lambda_C \cos^2(\alpha)}$$
(3)

137 where, λ_A , λ_B and λ_C are the three principal components of the thermal conductivity tensor; α , β and γ 138 are the angles between the line-source axis and the principal directions of thermal conductivity tensor 139 A, B and C, respectively (Fig. 1). Knowing the angles α , β and γ in equation (3), the apparent thermal 140 conductivity λ can be determined through three measurements.

141 Due to the stratification and bedding, many sedimentary rocks and clays are transversely isotropic, 142 leading to the following relations: $\lambda_A = \lambda_{\perp}$ (thermal conductivity in the direction perpendicular to the 143 bedding plane) and $\lambda_B = \lambda_C = \lambda_{\parallel}$ (thermal conductivity in the direction parallel to the bedding plane). If 144 the axis of the emitting needle probe is perpendicular to the bedding plane (Fig. 1a), $\alpha = 0^\circ$ and $\beta = \gamma$ 145 = 90°. Hence:

$$\lambda_0 = \sqrt{\lambda_B \lambda_C \cos^2(0^\circ)} = \lambda_{//} \tag{4}$$

146 When the needle is parallel to the bedding plane (Fig. 1b), $\alpha = \gamma = 90^\circ$, $\beta = 0^\circ$ and λ represents λ_{90} :

$$\lambda_{90} = \sqrt{\lambda_A \lambda_C \cos^2(0^\circ)} = \sqrt{\lambda_\perp \lambda_{//}}$$
(5)

147 From (4) and (5):

$$\lambda_{\perp} = \frac{\left(\lambda_{90}\right)^2}{\lambda_0} \tag{6}$$

Equation (6) has been used by several authors (Penner, 1963; Munroe and Sass, 1987; Popov et al., 149 1999; Pribnow et al., 2000; Gong, 2005; Davis et al., 2007; Popov et al., 2012; Riche and Schneebeli, 150 2012) to determine the value of thermal conductivity in the direction perpendicular to bedding λ_{\perp} 151 based on the measurements of λ_{90} and λ_0 in different materials (clay sediments, igneous rocks, snow, 152 etc.).

153

In general, when the line source axis forms an angle θ with the normal to the bedding (Fig. 1c), an apparent thermal conductivity can be determined by:

$$\lambda_{\theta} = \sqrt{\lambda_{\perp} \lambda_{//} \sin^2(\theta) + (\lambda_{//})^2 \cos^2(\theta)}$$
(7)

156 Thus,

$$\lambda_{\theta} = \sqrt{\left(\lambda_{0}\right)^{2} \cos^{2}(\theta) + \left(\lambda_{90}\right)^{2} \sin^{2}(\theta)} \tag{8}$$

Equation (8) was firstly proposed by Pribnow (cited by Davis et al., 2007). This function was used to
compare the experimental results with the changes in probe orientation. The data of Davis et al. (2007)
agree reasonably well with equation (8).

160 To calculate the value for $\theta = 45^{\circ}$, substituting equation (6) into equation (8) yields:

$$\lambda_{45} = \sqrt{\frac{\left[\left(\lambda_{0}\right)^{2} + \left(\lambda_{90}\right)^{2}\right]}{2}}$$
(9)

161

162 **2.2 Thermal conductivity at saturated state**

163 Various authors stated that the thermal conductivity of clays depends on several factors such as the 164 composition (solid, water and gas phases), the nature of the dominant minerals, the density/porosity, 165 the temperature, etc. (Farouki, 1986; Clauser and Huenges, 1995; Midttømme and Roaldset, 1999; 166 Tang et al., 2008). In this study, the effects of the slight changes in degree of saturation (S_r) and 167 density of the investigated samples were taken into account based on the Johansen's method (Farouki, 168 1986). According to Johansen, the thermal conductivity of an unsaturated soil is a function of its 169 thermal conductivity at dry and saturated states at the same dry density. To establish this function, 170 Johansen (1975) introduced a normalized thermal conductivity, namely Kersten's number (K_e):

$$K_e = (\lambda - \lambda_{dry}) / (\lambda_{sat} - \lambda_{dry})$$
⁽¹⁰⁾

171 where λ_{sat} is the thermal conductivity at saturated state, λ is the thermal conductivity at intermediate 172 degree of saturation, λ_{dry} is the thermal conductivity at dry state

Hence, one can deduce the saturated thermal conductivity value from the thermal conductivitymeasurement conducted on non-saturated samples, according to the following relationship:

$$\lambda_{sat} = \frac{(\lambda - \lambda_{dry})}{K_e} + \lambda_{dry}$$
(11)

175 In fine-grained unfrozen soils, Johansen proposes the following relationships for K_e and λ_{dry} :

$$K_{e} \cong \log S_{r} + 1.0 \tag{12}$$

176

$$\lambda_{dry} = \frac{0.135\rho_d + 64.7}{2700 - 0.947\rho_d} \pm 20\% \ (W/(m.K))$$
(13)

177 where ρ_d is the dry density in kg/m³.

178 **3 Material and methods**

179 **3.1 The Boom Clay**

180 In Belgium, researches on geological radioactive waste disposal in clay formation were initiated by 181 SCK+CEN in 1974. The first Underground Research Laboratory (URL) HADES was then excavated 182 in Boom Clay formation at a depth of 223 m during the early Eighties at a site close to the city of Mol 183 and has been progressively extended since. The Boom Clay was chosen as a potential host rock for 184 radioactive waste disposal due to its very low permeability (its hydraulic conductivity around 2.10⁻¹² m/s and 5.10^{-12} m/s in the direction perpendicular and parallel to the bedding, respectively), its 185 186 capacity to retard the migration of many radionuclides by sorption, its high plasticity, and its capacity 187 of self-sealing after physical disturbances.

188

Located at 190 m - 290 m below ground level at Mol (ONDRAF/NIRAS, 2001), Boom Clay was deposited during the Tertiary Era about 30 millions years ago. The Boom Clay layer is almost horizontal with a slight 1-2% north-east dip. It is located between water-bearing sand layers (ONDRAF/NIRAS, 2001).

193

194 In order to investigate both the anisotropy in thermal conductivity of Boom Clay and the possible 195 influence of the excavation damaged zone (EDZ) on it, nine samples were taken at various distances 196 from the gallery (4.0 m in diameter, 0.4 m in liner thickness) from the core (100 mm in diameter) 197 taken from a borehole (reference number: R66-67/2012) drilled horizontally till approximately 22 m in 198 length, in July 2012. Once drilled, the core was immediately sealed under vacuum in aluminium foils 199 to minimize drying effects. Table 1 summarizes the physical properties of the nine samples. The 200 distance to the gallery axis r varies from 2.5 m to 20.8 m. The water content of the samples (w)201 slightly changes from 22.6% to 24%, which is in the range of the in- situ measurements (from 22% to 27%). The densities are comparable (1630-1640 kg/m³) except for the sample that was the closest to 202 203 the gallery (TH6, r = 2.9 m, $\rho_d = 1.61$ Mg/m³). A suction measurement was made on sample TH11 (r= 9.60 m) using a chilled mirror hygrometer (WP4, Decagon Device), and a value of 2.8 MPa was 204

205 obtained, in good agreement with the suction value reported in Delage et al. (2007) from a sample at206 the same depth.

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208 **3.2** Thermal conductivity measurements

The apparatus used to measure the thermal conductivity is a commercial handheld device named KD2 Pro Thermal Properties Analyzer (Decagon Devices, Inc.). The kit consists of a controller and 3 separated sensors (small single needle, large single needle, dual needle) that can be inserted into the tested medium. The small single needle (60 mm in length, 1.3 mm in diameter) (Fig. 2b) was used to measure the thermal conductivity of Boom Clay. The accuracy of the thermal conductivity measurement using this needle is $\pm 5\%$ from 0.2 W/(m.K) - 2 W/(m.K).

215

216 Cylindrical samples were trimmed from cores (100 mm in diameter and 60mm - 90 mm in height).
217 The bedding orientation that could be identified from direct visual observation was marked. The
218 samples were then slightly confined by means of an adhesive tape so as to avoid further crack
219 propagation and any perturbation.

220

Three holes were prepared in each sample by drilled in each sample using a drilling-machine (Fig. 2a) equipped with a thin steel drill (1.3 mm in diameter, 65 mm in effective length). In order to ensure a good thermal contact between specimen and probe, the needle probe was coated with a thin layer of thermal grease (high-density polysynthetic silver thermal compound by Arctic Silver) prior to its insertion into the drilled hole (Fig. 2b), as recommended by the ASTM standard D5334-08 (2008). For a better precision, it was decided to run 5 successive measurements at intervals of 15 minutes.

227

The measured thermal conductivity values were denoted λ_0 , λ_{45} and λ_{90} , defined by the value of the angle θ between the axis of the hole and the direction normal to the bedding (see Fig. 1 and Fig. 3). As mentioned previously, heat transfer in the case of $\theta = 0^\circ$ is parallel to the bedding (Fig. 1a), and the λ_0 value measured by the needle device is equal to the thermal conductivity in the direction parallel to the bedding ($\lambda_{t/t}$). In the two other cases, the measured values depend on the thermal conductivities in the directions parallel and perpendicular to the bedding. The thermal conductivity in the direction perpendicular to the bedding (λ_{\perp}) needs to be back calculated, as seen in the previous section.

235

236 4 Experimental results

237 **4.1** Anisotropy of thermal conductivity

238 Since the degrees of saturation of the nine specimens are slightly lower than 100% (see Table 1), the 239 thermal conductivities at saturated state were first calculated using equation (11). As seen in Table 2 in which the initial $(\lambda_{l/0})$ and corrected $\lambda_{l/l}$ values are presented, the corrections made for saturation are 240 241 small with a relative variation $(\Delta \lambda_{ll} / \lambda_{ll0})$ included between 0.64% and 2.04%. The values of thermal 242 conductivity in the direction perpendicular to the bedding at saturated state λ_{\perp} were afterwards 243 determined using equation (6) based on the value of λ_{90} determined from a test with the needle parallel 244 to the bedding (Fig. 3). The thermal conductivity anisotropy ratio η was then deduced from equation 245 Erreur ! Source du renvoi introuvable..

246

247 **4.2 Influence of EDZ on thermal conductivity**

248 The changes in thermal conductivity along the three directions (λ_0 , λ_{45} , λ_{90} , Fig. 3) with respect to the 249 distance to the gallery axis r are shown in Fig. 4. The gallery wall is also represented in this figure. 250 The vertical bars indicate the range of the 5 measurements performed on each sample. The calculated 251 values of λ_{45} (from λ_0 and λ_{90} using equation (9)) are also plotted for comparison with the test data 252 directly obtained using the needle probe inclined at 45° with respect to the bedding. Excellent 253 agreement was observed, confirming: (i) 2D anisotropy model (presented in the part 2.1), (ii) 254 correspondence of principal axes of thermal conductivity to bedding plane orientation shown in Figure 255 1.

Fig. 4 shows that all the measurements start from a low value in the zone close to the gallery wall (around 1.3 W/(m.K) for $\lambda_0 = \lambda_{//}$ at r = 2.5 m). The conductivity values increase with distance to reach a maximum value at around 6.4 m from the gallery axis (1.6 W/(m.K) for $\lambda_{//}$). A slight decrease ($\lambda_{//}$) or a more significant one (λ_{\perp}) is observed for samples taken at greater distances from the gallery (up to 260 20.8 m). The average value of 1.6 W/(m.K) for $\lambda_{//}$ observed at larger distances is in good agreement 261 with the value of 1.65 W/(m.K) obtained from the back calculations carried out within the ATLAS III 262 project (Chen et al., 2011; Yu et al., 2011).

263

264 These results also confirm that the thermal conductivity in the direction parallel to the bedding is 265 always higher than that in the direction perpendicular to the bedding. The values of $\lambda_{//}$ vary between 266 1.34 - 1.61 W/(m.K) whereas those of λ_{\perp} vary between 0.82 – 1.06 W/(m.K) (Fig. 4). This can also be 267 observed in Fig. 5 that presents the changes in anisotropy ratios η versus the distance r to the gallery 268 axis. The values of η change between 1.5 and 1.8 with no significant effect of the distance to the gallery axis. Note that the value of λ_{\perp} obtained in this study is smaller than that obtained by back 269 270 calculations (1.31 W/(m.K)) by Chen et al. (2011). This difference may be explained by the fact that 271 λ_{\perp} is much more dependent on confinement (that strongly affects the contact between parallel clay 272 particles) than λ_{ll} . Further studies on the effect of confinement on the anisotropic thermal conductivity 273 are needed to confirm this point.

274

275 **5 Discussion**

276 It is now well recognized that an Excavation Damaged Zone (EDZ) characterized by a fracture 277 network including shear and tension fractures along with a herringbone pattern of fractures develops 278 during excavation in clays and claystones (Mertens et al., 2004; Bastiaens and Mertens, 2005; Davies 279 and Bernier, 2005; Tsang et al., 2005; Bernier et al., 2007; Tsang et al., 2012; Armand et al., 2013). 280 The extent of this zone depends on various parameters such as the nature of the media (soft clay, stiff 281 clay or claystone), its mechanical properties, the excavation technique, the state of stress around the 282 gallery and the bedding orientation. Around the Connecting Gallery of the URL HADES at Mol, an 283 EDZ characterized by a herringbone fracture pattern has been observed (Bastiaens and Mertens, 2005; 284 Bernier et al., 2007) with open fractures extending to about 1 m from the gallery wall. There were 285 indications that the extent in the horizontal direction was somewhat larger than in the vertical 286 direction.

288 The thermal conductivity measurements carried out in this study on samples extracted from a 289 horizontal borehole at various distances from the gallery axis evidenced changes that can be related to 290 the EDZ, given that discontinuities are suspected to affect heat transfer across these. In this regard, the 291 thermal conductivity changes in the direction parallel to the bedding (λ_{ij}) seems to be the more 292 significant parameter to be considered since the borehole was horizontal with a sub-vertical orientation 293 of cracks in this area. Indeed, inspection of Fig. 4 shows that parameter λ_{ij} starts from small values to 294 afterwards increase with the distance to the gallery (until a distance r of 4.0 m from the gallery axis, 295 corresponding to about 1.6 m from the outside wall). Beyond this distance, the horizontal thermal 296 conductivity remains about constant, around a value of 1.6 W/(m.K), in good accordance with the 297 back analysis of the in-situ heating test. Based on the thermal conductivity measurements, it is 298 deduced that the EDZ extension in the horizontal direction is about 1.6 m from the outside of gallery's 299 wall, a value in consistence with in situ observation.

300

301 The data presented here were obtained from the laboratory measurements carried out on samples 302 extracted from a horizontal core and under zero confining stress. As a result, there would be some 303 possibility of cracks' opening within the samples due to stress release. Thereby, the effects of cracks 304 on thermal fluxes observed in Fig. 4 could be enhanced by the stress release. Thus, the question arises 305 as to whether the changes in thermal conductivity with respect to the distance to the gallery really 306 occur in-situ. Actually, clays are known to exhibit self-sealing properties with respect to water 307 transfer. This self-sealing behaviour, interpreted as the recovery of a good contact between the two 308 faces of discontinuities, could have comparable consequence on heat transfer as well. In other words, 309 self-sealing could be extended from fluid transfer to heat transfer. Such a hypothesis should obviously 310 be further investigated by thermal conductivity measurements in the laboratory, under confining 311 conditions and/or running detailed in-situ measurement at various distances from the gallery 312 completed by thermal back-analysis of in-situ thermal tests. Meanwhile, the changes in thermal 313 conductivity carried out here on laboratory samples provide a complementary identification of the 314 extent of the EDZ along a horizontal direction around the gallery excavated in Boom Clay.

316 6 Conclusion

317 An experimental study using the needle thermal probe technique was conducted for investigating the 318 thermal conductivity anisotropy of natural Boom Clay. The investigation was carried out on samples 319 extracted from a series of cores (from a single borehole) taken from the Underground Research 320 Laboratory at Mol in Belgium, at a depth of 223 m. Thermal conductivity was measured on specimens 321 sampled at various distances to the gallery axis, from 2.5 to 20.8 m. Measurements were conducted 322 along three directions: parallel, perpendicular and at an angle of 45° with respect to the bedding plane. 323 A significant thermal anisotropy was evidenced and data confirmed that the highest conductivity was 324 in the direction parallel to bedding and that the smallest one was in the direction perpendicular to 325 bedding, the one at 45° being logically in between these two values. It is worth noting that the 326 conductivity values found in this laboratory study in the direction parallel to the bedding was close to 327 that back-calculated from the in-situ ATLAS experiments (Chen et al., 2011).

The effect of an excavation damaged zone (EDZ) was also evidenced: the thermal conductivity in the three directions increased with the distance to the gallery. Based on the experimental data, an EDZ extent of about 1.6 m in the horizontal direction was identified, a value consistent with the in situ observation. Finally, the thermal anisotropy did not appear to be significantly affected by the excavation damage, with no significant effect of the distance to the gallery on the thermal anisotropy ratio that was comprised between 1.48 and 1.74.

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 Repository for Heat-Emitting High-Level Radioactive Waste in a Clay Formation: a Performance
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459 List of Tables

- 460 Table 1: Properties of natural Boom Clay samples taken at various distances togallery axis (ρ_d : dry
- 461 density, w: water content, e: void ratio, S_r : degree of saturation).
- 462 Table 2: Initial and corrected $\lambda_{//}$ values ($\lambda_{//0}$: thermal conductivity measured on non-saturated samples,
- 463 Corrected $\lambda_{l/l}$: saturated thermal conductivity after Johansen's method, $\Delta \lambda_{l/l}$: difference between the
- 464 $\lambda_{1/0}$ and Corrected $\lambda_{1/1}$).
- 465

466 List of Figures

- 467 Fig.1: Principal axes of thermal conductivity for the three cases studied with the needle probe method.
- 468 A, B and C: main (orthogonal) axes of thermal conductivity (A is perpendicular to the bedding plane;
- 469 B and C are parallel to the bedding plane).
- 470 Fig.2: Running a test by the needle probe method (a): system for drilling; (b) inserting the needle
- 471 probe into the specimen with marks of the bedding plane.
- 472 Fig.3: Measurement of thermal conductivity.
- 473 Fig.4: Thermal conductivity versus distance *r* to the axis of the gallery.
- 474 Fig.5: Degree of thermal conductivity anisotropy versus distance *r* to the axis of the gallery.
- 475

Test	Distance r (m)	$\rho_{\rm d}~(\rm kg/m^3)$	w (%)	е	S _r (%)
TH6	2.5	1610	23.98	0.65	97.9
TH7	2.6	1640	23.05	0.63	98.0
TH8	2.7	1640	23.05	0.63	98.0
TH14	3.4	1630	23.51	0.64	98.4
TH13	3.8	1630	23.02	0.63	98.0
TH12	6.0	1630	22.84	0.64	95.2
TH11	9.2	1640	22.63	0.63	96.3
TH10	16.0	1640	22.61	0.63	96.2
TH9	20.8	1630	23.48	0.64	98.4

477 Table 1: Properties of natural Boom Clay samples taken at various distances *r* to gallery axis (ρ_d : dry

density, w: water content, e: void ratio, S_r : degree of saturation)

483 Table 2: Initial and corrected $\lambda_{//}$ values ($\lambda_{//0}$: thermal conductivity measured on non-saturated samples, 484 Corrected $\lambda_{//}$: saturated thermal conductivity after Johansen's method, $\Delta \lambda_{//}$: difference between the 485 $\lambda_{//0}$ and Corrected $\lambda_{//}$).

Tosts	Distance <i>r</i>	S	λ//0	Corrected $\lambda_{\prime\prime}$	$\Delta \lambda_{//} / \lambda_{//measured}$)
16818	(m)	S_r	(W/(m.K))	(W/(m.K))	(%)
TH6	2.5	97.9	1.33	1.34	0.85
TH7	2.6	98.0	1.38	1.39	0.80
TH8	2.7	98.0	1.41	1.42	0.80
TH14	3.4	98.4	1.45	1.46	0.64
TH13	3.8	98.0	1.54	1.55	0.81
TH12	6.0	95.2	1.56	1.59	2.04
TH11	9.2	96.3	1.58	1.61	1.55
TH10	16.0	96.2	1.49	1.51	1.57
TH9	20.8	98.4	1.52	1.53	0.67



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491 Fig. 1: Principal axes of thermal conductivity for the three cases studied with the needle probe method.
492 A, B and C: main (orthogonal) axes of thermal conductivity (A is perpendicular to the bedding plane;

493 B and C are parallel to the bedding plane).

494



495

- 496 (a) (b)
- 497 Fig. 2 : Running a test by the needle probe method (a): system for drilling; (b) inserting the needle
- 498 probe into the specimen with marks of the bedding plane.

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488







504 Fig. 4: Thermal conductivity versus distance *r* to the axis of the gallery.



Fig. 5: Degree of thermal conductivity anisotropy versus distance r to the axis of the gallery.